

# The interior structure of Mercury constrained by geodesy data

Attilio Rivoldini

# New Geodesy data

- Margot et al. 2012  
Obliquity:  $2.04 \pm 0.08$  arcmin  $\Rightarrow$   $\text{MOI} = 0.345 \pm 0.014$  (4%)  
Libration amplitude:  $38.5 \pm 1.6$  arcsec (4%)

- Stark et al. 2015  
Obliquity:  $2.03 \pm 0.09$  arcmin  $\Rightarrow$   $\text{MOI} = 0.345 \pm 0.014$   
Libration amplitude:  $38.9 \pm 1.3$  arcsec

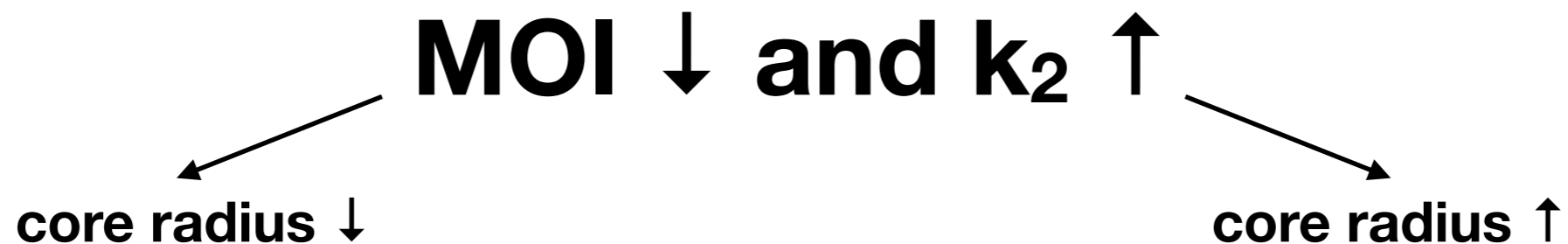
- Verma et al. 2016  
Tidal Love number:  $k_2 = 0.46 \pm 0.02$

- Genova et al. 2019  
Obliquity:  $1.97 \pm 0.009$  arcmin  $\Rightarrow$   $\text{MOI} = 0.333 \pm 0.0015$  (0.5%)  
Libration amplitude:  $40.0 \pm 8.7$  arcsec  
Tidal Love number:  $k_2 = 0.57 \pm 0.03$  (5.2%)

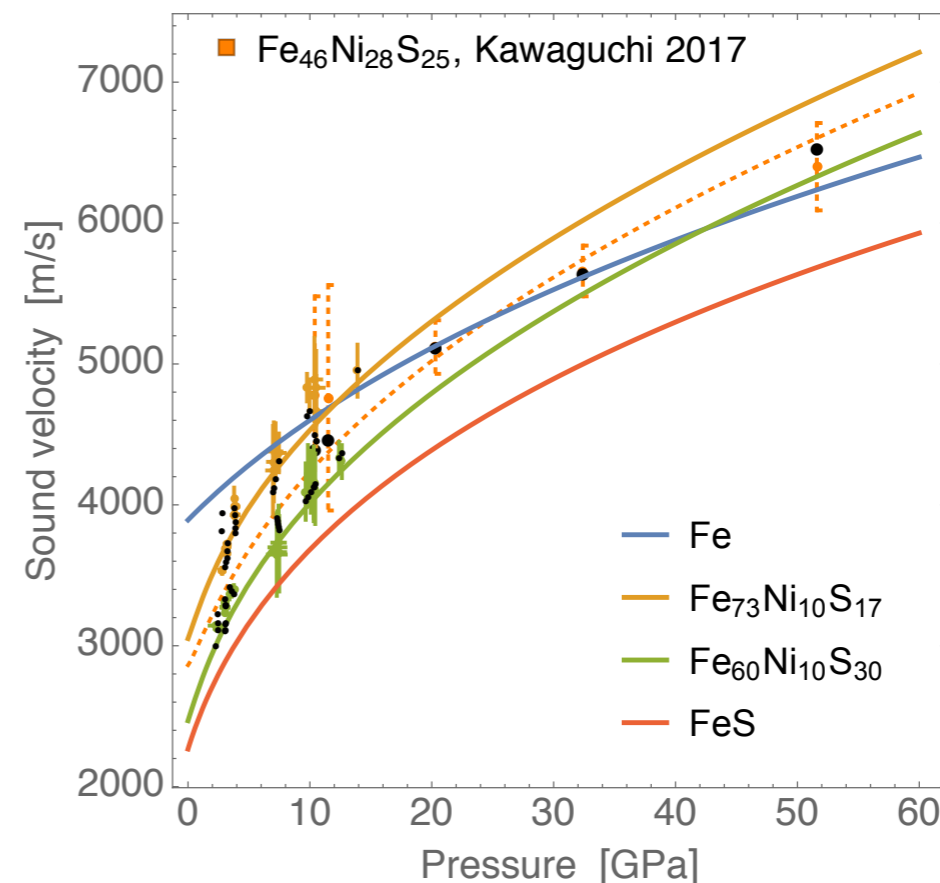
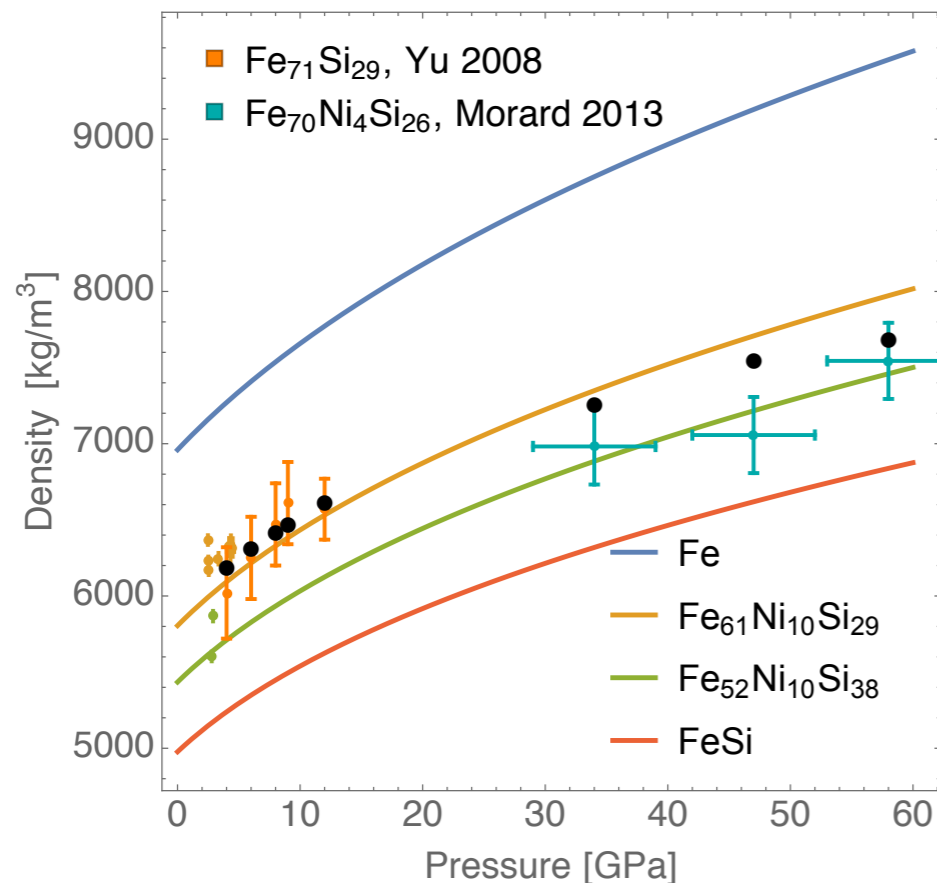
- Konopliv et al. 2020  
Obliquity:  $1.99 \pm 0.12$  arcmin  $\Rightarrow$   $\text{MOI} = 0.337 \pm 0.02$  (~6%)  
Tidal Love number:  $k_2 = 0.53 \pm 0.03$  (5.6%)

OLD

NEW



# New thermoelastic data about I-Fe alloys



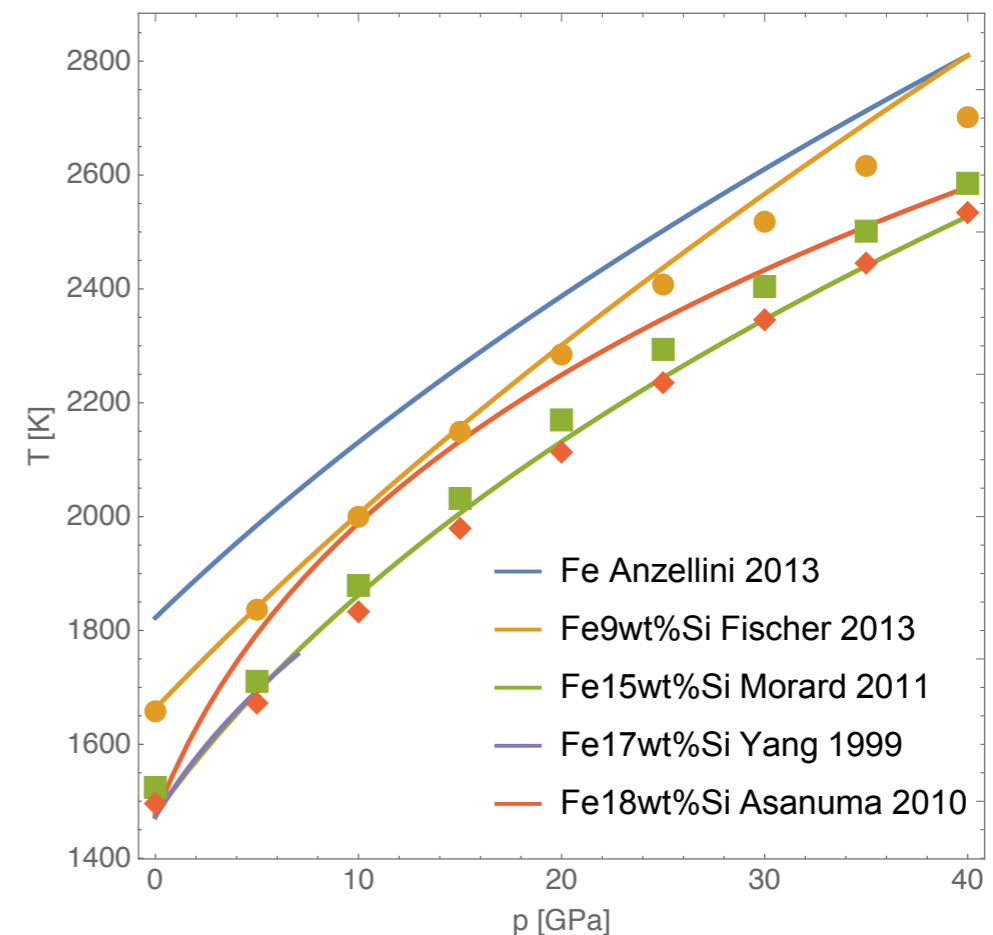
- Terasaki et. al 2019: densities (up to 5 GPa) and acoustic sound velocities (up to 14 GPa) of (Fe<sub>73</sub>Ni<sub>10</sub>S<sub>17</sub>, Fe<sub>60</sub>Ni<sub>10</sub>S<sub>30</sub>) and (Fe<sub>61</sub>Ni<sub>10</sub>Si<sub>29</sub>, Fe<sub>52</sub>Ni<sub>10</sub>Si<sub>38</sub>)
- non-ideal solution model required to summarize Fe-S and Fe-Si elastic data
- liquid solution model in agreement with low pressure and high pressure data

# Prior assumptions, modeling, and data

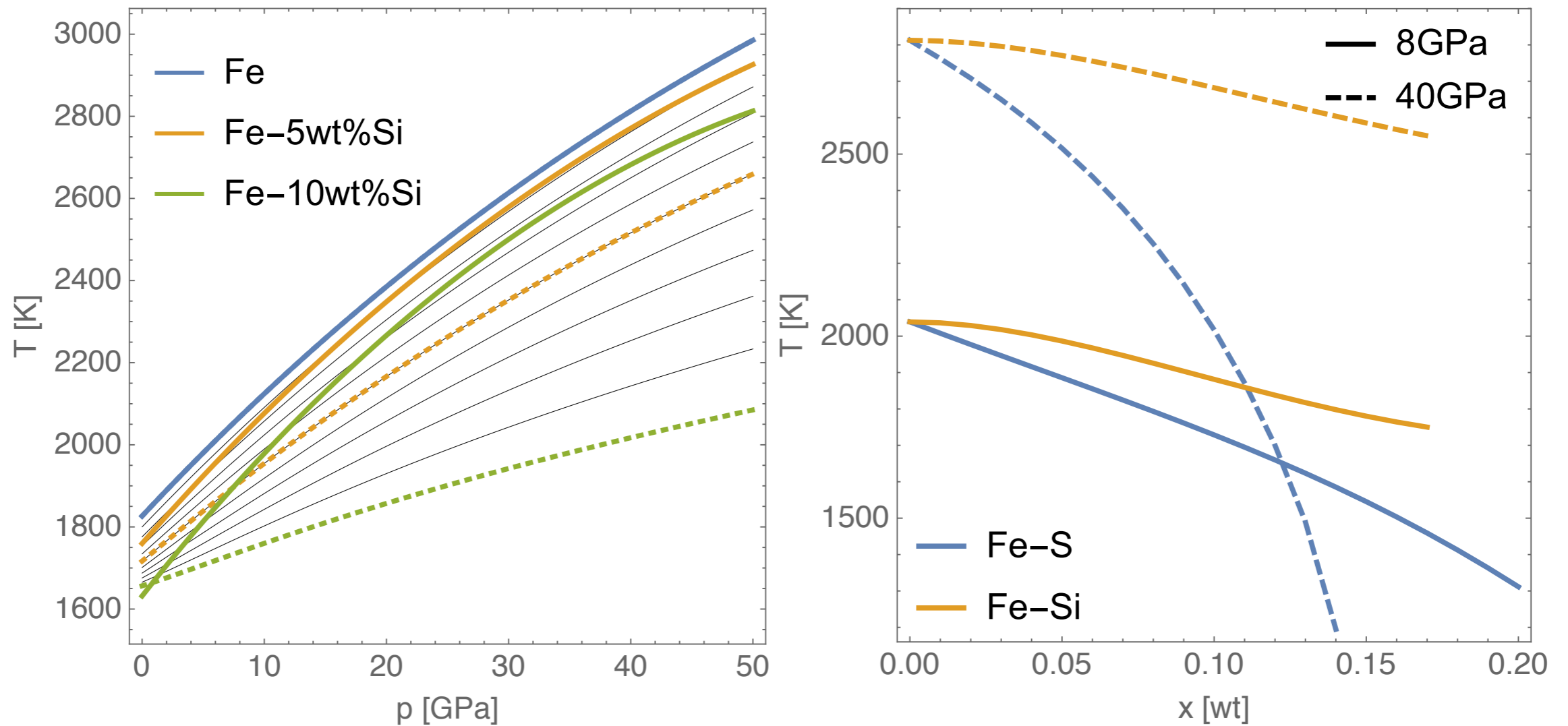
- crust: density [2700, 3100]kg/m<sup>3</sup> and thickness [15,120]km
- mantle elastic properties compatible with forsterite-enstatite mixture; corresponds to a prior mantle density [3150, 3400]kg/m<sup>3</sup>
- inner core radius and light element fraction in agreement with liquidus; assume Si concentration in liquid below eutectic composition (because of unknown liquidus at those compositions); core radius prior [1800, 2200]km
- prior core-mantle boundary temperature between eutectic temperature and (very likely too warm!) mantle solidus
- libration amplitude calculated by taking into account gravitational core-mantle coupling and mantle induced core density stratification (Dumberry et al., 2013)
- geodesy data: MOI and  $k_2$  from Konopliv 2020 and libration amplitude Margot et al. 2012

# Fe-Si melting

- based on 1 bar Fe-rich liquidus, and melting data of: Fe9wt%Si, Fe15wt%Si, Fe15wt%Si, Fe18wt%Si
- assume equipartitioning of Si in solid and liquid Fe
- Margules model with pressure dependent interaction coefficients

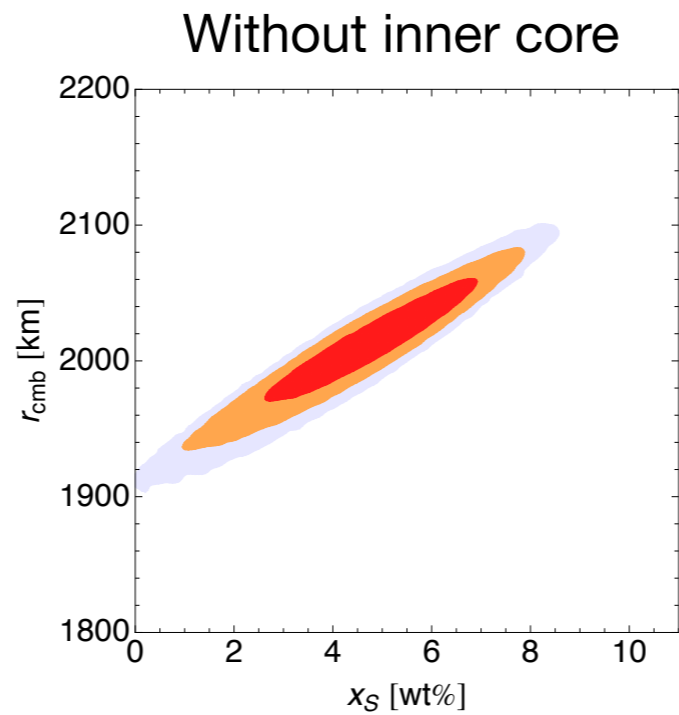
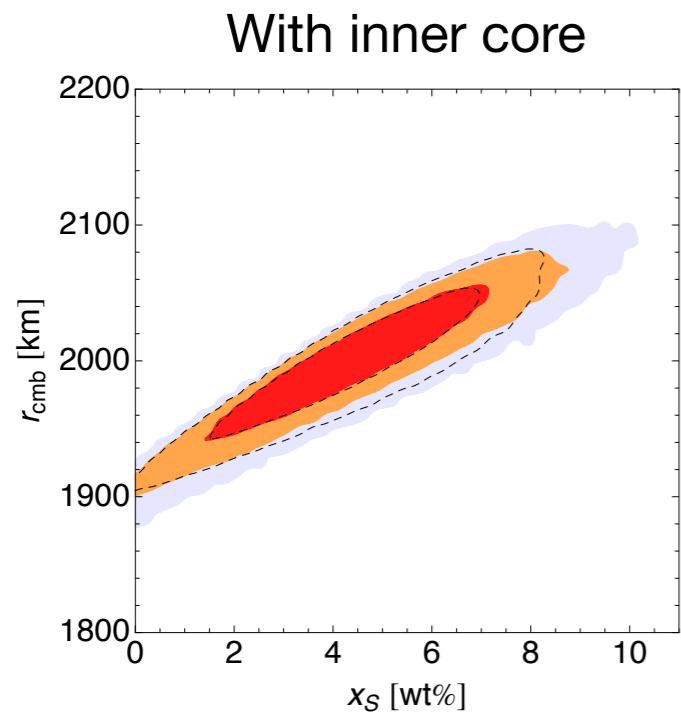


# Fe-Si melting compared to Fe-S



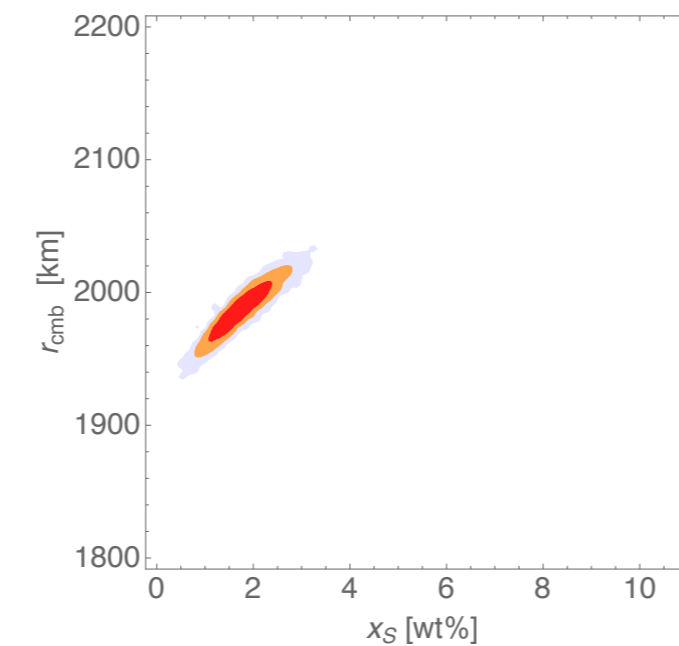
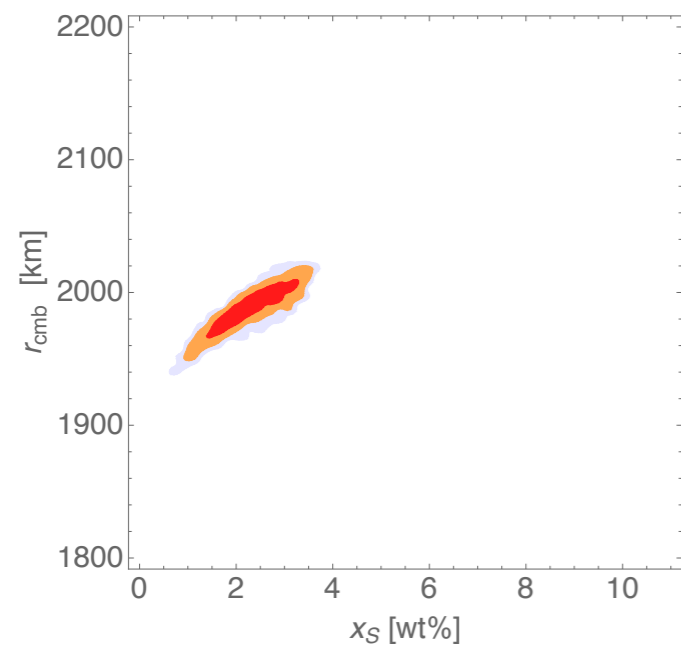
⇒ 5wt%Si decrease Fe liquidus by about as much as 1wt%S

# Results: Fe-S models compared to Rivoldini et al. 2013



2013

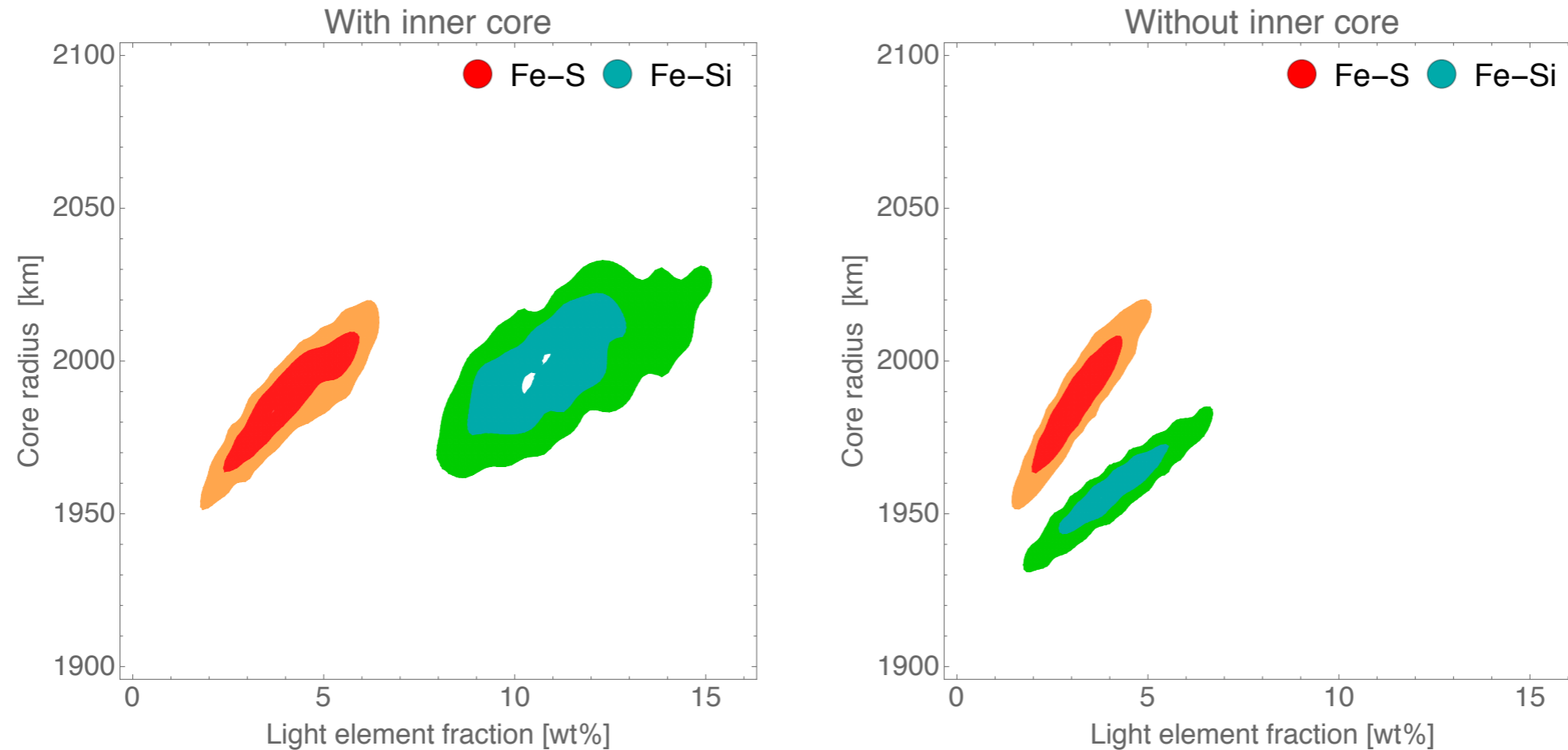
	2013	2020
MOI	$0.346 \pm 0.014$	$0.337 \pm 0.02$
$g_{88}$	$38.5'' \pm 1.6''$	$38.5'' \pm 1.6''$
$C_m$	$0.148 \pm 0.006$	$0.148 \pm 0.006$
$k_2$	-	$0.53 \pm 0.03$



2020

	2013	2020
$r_{cmb}$ [km]	1965-2043	1974-2003
$r_{icb}$ [km]	345-1430	246-899
$x_S$ [wt%]	2.8-6.2	2.7-5.3
$\rho_{Mantle}$ [kg/m <sup>3</sup> ]	3163-3449	3110-3180

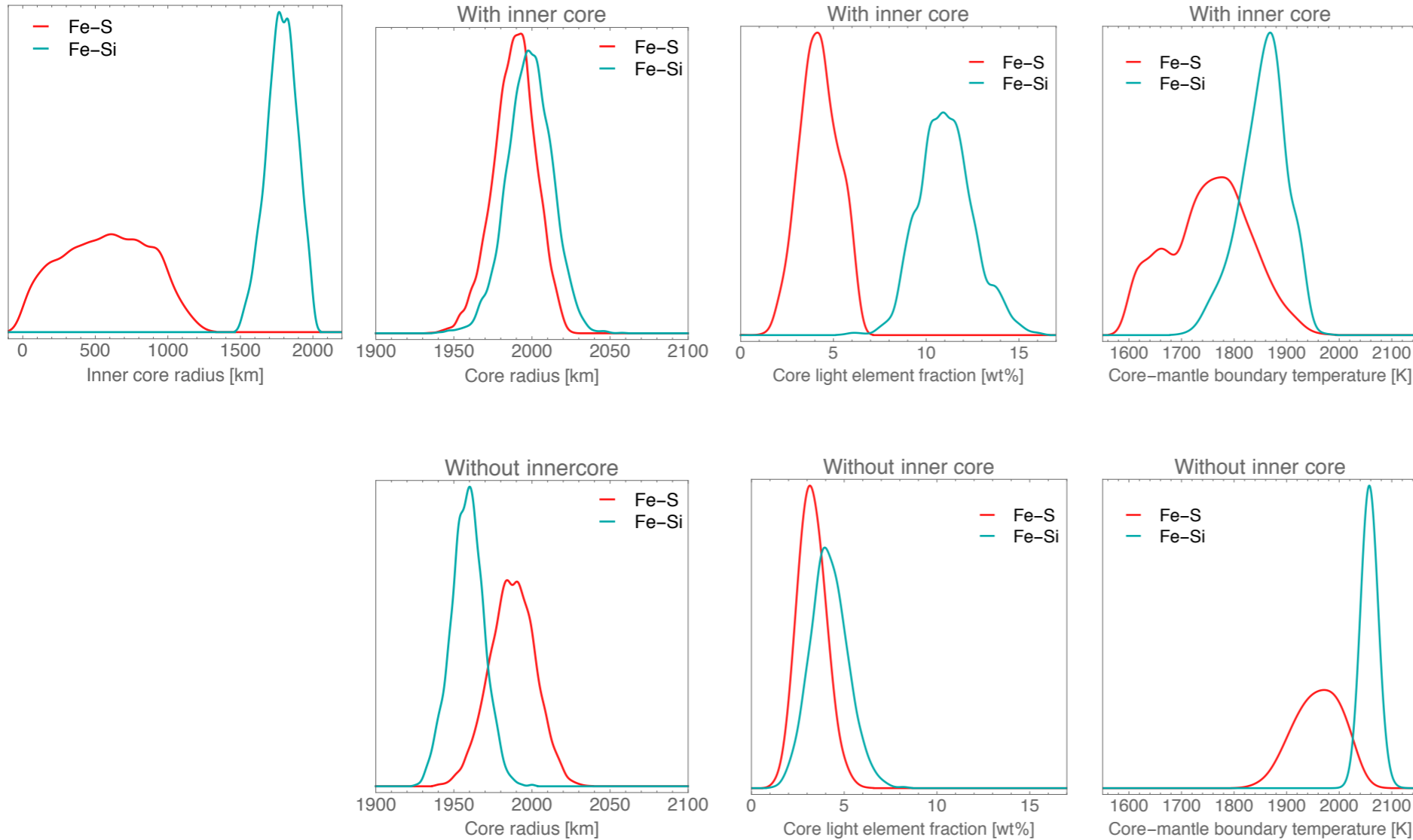
# Results



- core radius between ~1926 and 2030km
- Fe-Si models require more light elements since Fe-Si alloys are denser than Fe-S alloys



# Results



## With inner core

	Fe-S	Fe-Si
$r_{\text{cmb}}$ [km]	1975-2002	1984-2012
$r_{\text{icb}}$ [km]	236-896	1673-1892
$x_{\text{S}}$ [wt%]	3.2-5.3	9.6-12.5
$T_{\text{cmb}}$ [K]	1664-1831	1814-1892

## Without inner core

	Fe-S	Fe-Si
$r_{\text{cmb}}$ [km]	1973-2002	1948-1968
$x_{\text{S}}$ [wt%]	2.5-4.0	3.2-5.2
$T_{\text{cmb}}$ [K]	1913-2006	2051-2064

- models with and without inner core have a comparable core radius range with the exception of the liquid Fe-Si models
- liquid Fe-Si models only possible for high  $T_{\text{cmb}} \geq 2037\text{K}$  ( $3\sigma$ ) (implying molten lower mantle?) and inner core models with Fe-Si only possible if  $T_{\text{cmb}} \geq 1718\text{K}$  ( $3\sigma$ )
- Fe-S models  $r_{\text{icb}} \sim [0, 1214]\text{km}$  and Fe-Si  $r_{\text{icb}} \sim [1507, 2000]\text{ km}$  ( $3\sigma$ )
- mantle density  $\sim [3100, 3310]\text{kg/m}^3$  ( $3\sigma$ )
- almost no constraints on crust thickness and density

# Conclusions

- models with and without an inner core agree with the new geodesy data (liquid Fe-Si models only possible if  $T_{\text{cmb}} \gtrsim 2037\text{K}$ )
- new smaller MOI value favors models with smaller core radii Fe-S [1975,2003]km (old [1965,2043]km) and FeSi [1950-2012]km as well as a smaller mantle density ( $< 3310 \text{ kg/m}^3$ )
- Fe-S models have an inner core radius below about 1300km and for Fe-Si models the inner core radius is  $\sim [1500,2000]\text{km}$
- new MOI value favors models with smaller cores but those models are somewhat at odds with the large  $k_2$  value suggesting that the mantle in the models is too stiff